SEDI 2020 Session 7 – Inner Core

Jenny Wong ISTerre, Grenoble









Natural Environment Research Council



Introduction

- Present day radius is 1220 km
- First crystallised approximately 1 billion years ago (debated)
- Composed of solid iron containing light element impurities
- Continuously freezing and growing as the planet cools over time
- Latent heat release and partitioning of light elements into the liquid core is an important power source for the geodynamo

Seismic observations

Year	Event	10.4 (km/s)
1953	Bullen alphabetises the structure of the Earth and names "shell F"	10.3
1972-73	Cleary & Haddon, King et al. debunk the early F-layer	
1981	Dziewonski et al. publish PREM	10.2 - PREM AK135 S&P
1991	Souriau & Poupinet detect the modern F-layer with $d = 150 \ km$	10.1 SP6 VMOI OW SPR 1200 1
onwards	Many studies support this observation with $d = 150$ to $400 \ km$	



Density structure

- Detectable by Slichter modes?
- Also have a discrepancy between $\Delta \rho_{mod}$ and $\Delta \rho_{bod}$
- This infers that a stably stratified layer indeed exists
- How can light elements pass through the F-layer and out into the bulk of the liquid core?
- Layer cannot be a thermal boundary layer



Possible scenarios







Thermochemical layer on the liquidus

Convective translation

Slurry layer

Model details



- Two component (iron and oxygen), two phase (solid and liquid) system
- Formation and transport of solid phase provides a way for light elements to pass through a stablystratified layer
- Solid fraction is small

For full details please see Wong *et al.* (2018) <u>https://doi.org/10.1093/gji/ggy245</u>

Model details



- Governing equations:
 - Liquidus
 - Conservation of oxygen
 - Conservation of energy
- Steady slurry, 1D and spherical geometry
- Reference frame of fixed layer thickness moving at IC growth rate

Geophysical constraints

	$\Delta ho_{mod} \left(kg \ m^{-3} ight)$	$\Delta ho_{bod} \left(kg \ m^{-3} ight)$	CMB heat flow (<i>TW</i>)	ICB heat flow (TW)
Maximum	1000 (Masters & Gubbins 2003)	1100 (Tkalčić et al. 2009)	15 (Lay et al. 2008)	2 (Pozzo et al. 2014)
Minimum	600 (PREM)	520 ± 240 (Koper and Dombrovskya 2005)	5 (Lay et al. 2008)	> 0

Model solution

- Temperature gradient is "locked" to the oxygen gradient via the liquidus
- Solid flux is negative downward towards ICB
- Increasing layer thickness destabilises the layer



Model solution

- Temperature gradient is "locked" to the oxygen gradient via the liquidus
- Solid flux is negative downward towards ICB
- Increasing layer thickness destabilises the layer
- Temperature, oxygen and solid flux contribute to density anomaly



Regime diagram

- Perform parameter search over *Pe,St,Le* space (fixed layer thickness)
- Determine which parameter combinations give
 - a **STABLE** slurry
 - a PARTIALLY STABLE slurry
 - an UNSTABLE slurry
 - NO slurry
- Apply geophysical constraints to narrow solution space
 - Density jump $(\rho_s \rho_{sl})$
 - CMB heat flow



Geophysical implications



$$v_p^2 = \frac{K}{\rho}$$

$$d = 250 \ km$$
$$k = 100 \ W \ m^{-1} K^{-1}$$
$$Q_i = 3.5 \ TW$$
$$Q_{sl} = 6 \ TW$$

Geophysical implications





Motivation

Consensus on slowdown in P-wave speed at the base of the core.



Slurry (iron snow) layer

A slurry provides an explanation of how the stably-stratified F-layer is maintained.



Regime diagram

Stable slurry is likely when $Pe \gtrsim Le$. Agrees with available geophysical constraints.



Ongoing work

How does a stably-stratified F-layer impact core dynamics and the dynamo?

References:

- Alboussière et al. (2010). Melting-induced stratification above the Earth's inner core due to convective translation, Nature, 466. Aubert (2019). Approaching Earth's core conditions in high-resolution geodynamo simulations, GJI, 219(S1). Bullen (1954). An introduction to the theory of seismology, University Press. Gubbins et al. (2008). A thermochemical boundary layer at the base of Earth's outer core and independent estimate of core heat flux, GJI, **174.** Jacobs (1953). The Earth's inner core. Nature, 172 King et al. (1973). Evidence for seismic wave scattering in the D" layer , EPSL, 20. Labrosse (2015). Thermal evolution of the core with a high thermal conductivity, PEPI, 247. Loper & Roberts (1977). On the motion of an iron-alloy core containing a slurry: I. General theory, GAFD, 9. Ohtaki et al. (2015). Independent estimate of velocity structure of Earth's lowermost outer core beneath the northeast Pacific from PKiKP-PKPbc differential traveltime and dispersion in PKPbc, JGR: Solid Earth, **120.** Ohtaki et al. (2018). Seismological evidence for laterally heterogeneous lowermost outer core of the Earth, JGR: Solid Earth, 123. Rosat et al. (2006). A search for the Slichter triplet with superconducting gravimeters: Impact of the density jump at the inner core boundary, JoGeodynamics, **41**.
- Slichter (1961). The fundamental free mode of the Earth's inner core, *PNAS*, **47**. Williams (2018). The thermal conductivity of Earth's core: A key physical parameter's
- constraints and uncertainties, *Ann. Rev. Earth and Planetary Sciences,* **XV.** Wong et al. (2018). A Boussinesq slurry model of the F-layer at the base of Earth's outer core, *GJI*, **214.**

Special thanks to:

Séverine Rosat Chris Davies Chris Jones Julien Aubert Renaud Deguen Université de Strasbourg University of Leeds University of Leeds IPGP ISTerre



@

jenny.wong@univ-grenoble-alpes.fr

🄰 @_jennywong_



<u>www.jnywong.github.io</u>







Engineering and Physical Sciences Research Council



